Evaluation of Planetary Boundary Layer Schemes in Hurricanes Over Land Through Comparison of Surface Winds in Observations and Simulations of Hurricane Wilma (2005)

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Motivation

- Global and regional dynamical models are the primary tools for predicting the impacts of hurricanes.
- Sut, these are not used to predict the wind fields over land.
- Probably for two good reasons:
 - Typical forecast track and intensity errors could cause huge errors in local wind forecasts.
 - Forecast model over-land winds have not been validated, or improved.



Previous Work

 Nolan et al. (2009a,b) performed a comprehensive evaluation of the Yonsei University (YSU) and Mellor-Yamada-Janjic (MYJ) boundary layer parameterizations in WRF simulations of Hurricane Isabel (2003) against point observations and synthesized analyses.



MYJ

 $\frac{d}{dt}\left(\frac{q^2}{2}\right) - \frac{\partial}{\partial z}\left[lqS_q\frac{\partial}{\partial z}\left(\frac{q^2}{2}\right)\right] = P_s + P_b - \varepsilon \quad \text{(TKE)}$ $l = l_0\kappa z(\kappa z + l_0)^{-1} \text{ (geometrically prescribed length scale varies from 0 to } l_0\text{)}$ $K_M = lqS_M \text{, and } K_T = lqS_T \quad \text{(eddy diffusivities of momentum and temperature)}$

Through fundamentally different approaches, both schemes reproduced the hurricane boundary layer quite well. YSU was a little better, as the MYJ caused too much loss of angular momentum causing an "exaggerated" secondary circulation.

What about over land?

Modeling Parameters & Strategy

• WRF 3.9.1.1

- 9 km / 3 km / 1 km domains
- 60 vertical levels between 42 m 20 km
- YSU and MYJ boundary layer schemes
 - MYJ surface drag formula over water changed to match YSU formula
- GFS initial and boundary conditions
 - Domain-scale grid nudging to GFS analyses at 1/12h timescale.
- Surface winds on 1 km nest saved every 10 seconds
- Simulation begins ~28 h before landfall
- Interpolation of high-resolution terrain & land use data to 3 km and 1 km nested grids as they move (Chen et al. 2007)



85°W

90°W

Roughness Length z_0 in d03

WRF 3.9.1.1 Domain Configuration, GFS Total Wind Speed at z=1.43k



Vortex Bogussing

Vortex "bogussing" method of Rappin et al. (2013) is used to replace the GFS vortex with a smaller, stronger vortex in a better location.

- Vortex removal technique modeled after GFDL system.
- Vertical structure of wind field built from MPI theory of Emanuel (1986).
- Trial-and-error to make a simulation with track, intensity, and size as close as possible to Wilma (2005).



Track, Intensity, and Size



GFS initial conditions produced excellent track & intensity.
But the vortex size was too large.





YSU





RADIUS FROM CENTER (km)

MYJ





4.2 AMS Tropical Meteorology and Tropical Cyclones Symposium, 100th Annual Meeting

Results & Comparisons



Surface winds reduced over land in both models, but more in YSU

- YSU winds further reduced in urban areas; seems more correct
- Wind streaks are associated with mesoscale vortices 5-10 km in scale

Results & Comparisons

 The real test: Comparison to time series from point observations.

 Data from airports (MIA, FLL, PBI) and from mobile towers deployed by the *Florida Coastal Monitoring Program* (Masters et al. (2010), Balderrama et al. (2011))





Model winds are 2-minute means, like ASOS at airports.

- MYJ is nearly perfect! YSU is much less. Why?
- The land use data set has the airport location designated as "urban". But airport measurements are often in "open exposure".



• YSU time series corrected to $Z_{open} = 0.03$ m using the logarithmic law:

• $S10_{open} = S10 \times \frac{\log(Z_1/Z_0)}{\log(10/Z_0)} \times \frac{\log(10/Z_{open})}{\log(Z_1/Z_{open})}$ where Z_l = lowest model level

 MYJ does not seem to recognize rougher surfaces, so this correction causes overestimates...do not apply it.

Wind Speed

10m

Similar agreement from both schemes at Miami (top) and Fort Lauderdale (bottom) airports





- 10 Hz FCMP tower data converted to 3-s gusts, 1-min winds, and 10-min winds.
- Model cannot reproduce the variability of the 1-min winds.
- But it can reproduce the variability of 10-min winds.



 Similar agreement from both schemes at Tower2 (top), but a bit high at Tower3 (bottom)

Miami's FIU campus...
 z₀ > 0.5m?

4.2 AMS Tropical Meteorology







Normalized azimuthal-mean vertical profiles of wind over land and water.

- Nearly identical over water (left)
- Over land, YSU has reduced surface winds (upper right)
- Over land, MYJ has reduced winds in lowest model level (lower right)



- Compare vertical wind profiles in hurricane conditions over "urban" area...
 - 2-8% different at model levels in lowest 500 m
 - About 30% different at the surface

• WHY?



dqx = min(qx(i,k+1)-qx(i,k),0.0)hfxpbl(i) = we(i)*dthx qfxpbl(i) = we(i)*dqx

/SU dux = ux(i, k+1) - ux(i, k)dvx = vx(i,k+1)-vx(i,k)if(dux.gt.tmin) then ufxpbl(i) = max(prpbl(i)*we(i)*dux,-ust(i)*ust(i)) elseif(dux.lt.-tmin) then ufxpbl(i) = min(prpbl(i)*we(i)*dux,ust(i)*ust(i)) else ufxpbl(i) = 0.0endif if(dvx.gt.tmin) then vfxpbl(i) = max(prpbl(i)*we(i)*dvx,-ust(i)*ust(i)) elseif(dvx.lt.-tmin) ther vfxpbl(i) = min(prpbl(i)*we(i)*dvx,ust(i)*ust(i)) else vfxpbl(i) = 0.0 endif delb = govrth(i)*d3*hpbl(i) delta(i) = min(d1*hpbl(i) + d2*wm2(i)/delb,100.) endif enddo do k = kts,klpbl do i = its,ite if(pblflg(i).and.k.ge.kpbl(i))then entfac(i,k) = ((zq(i,k+1)-hpbl(i))/delta(i))**2. else entfac(i,k) = 1.e30endif enddo enddo compute diffusion coefficients below pb] do k = kts,klpbl do i = its,ite if(k.lt.kpbl(i)) then zfac(i,k) = min(max((1.-(zq(i,k+1)-zl1(i))/(hpbl(i)-zl1(i))),zfmin),1.) z facent(i,k) = (1,-z fac(i,k)) **3.wscalek(i,k) = (ust3(i)+phifac*karman*wstar3(i)*(1,-zfac(i,k)))**h1wscalek2(i,k) = (phifac*karman*wstar3_2(i)*(zfac(i,k)))**h1 if(sfcflg(i)) then prfac = conpr prfac2 = 15.9*(wstar3(i)+wstar3_2(i))/ust3(i)/(1.+4.*karman*(wstar3(i)+wstar3_2(i))/ust3(i)) prnumfac = $-3.*(\max(zq(i,k+1)-sfcfrac*hpbl(i),0,))**2./hpbl(i)**2.$ else prfac = 0. prfac2 = 0.prnumfac = 0. phim8z = 1.+aphi5*zol1(i)*zq(i,k+1)/zl1(i) wscalek(i,k) = ust(i)/phim8z wscalek(i,k) = max(wscalek(i,k),0.001) endif prnum0 = (phih(i)/phim(i)+prfac) prnum0 = max(min(prnum0,prmax),prmin) xkzm(i,k) = wscalek(i,k) *karman* zq(i,k+1) * zfac(i,k)**pfac+ & wscalek2(i,k)*karman*(hpbl(i)-zq(i,k+1))*(1-zfac(i,k))**pfac !Do not include xkzm at kpbl-1 since it changes entrainment if (k.eq.kpbl(i)-1.and.cloudflg(i).and.we(i).lt.0.0) then xkzm(i,k) = 0.0endif prnum = 1. + (prnum0-1.)*exp(prnumfac) xkzq(i,k) = xkzm(i,k)/prnum*zfac(i,k)**(pfac_q-pfac) prnum0 = prnum0/(1.+prfac2*karman*sfcfrac) prnum = 1. + (prnum0-1.)*exp(prnumfac) xkzh(i,k) = xkzm(i,k)/prnumxkzm(i,k) = xkzm(i,k)+xkzom(i,k)xkzh(i,k) = xkzh(i,k)+xkzoh(i,k)xkzq(i,k) = xkzq(i,k)+xkzoh(i,k)xkzm(i,k) = min(xkzm(i,k),xkzmax)xkzh(i,k) = min(xkzh(i,k),xkzmax)xkzq(i,k) = min(xkzq(i,k),xkzmax) endif

REAL, DIMENSION (MSS: MSE, KTS: KTE), INTENT (INOUT):: SPECIES

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!***
!*** LOCAL VARIABLES
!***
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INTEGER:: K,M

REAL:: CF, CMB, CMSB, DTOZL, DTOZS, RCML, RHOK, RKHH, RKHZ, RKSS, RSSB

REAL, DIMENSION(KTS:KTE-1):: CM, CR, DTOZ

REAL, DIMENSION (MSS: MSE, KTS: KTE-1):: RKCT, RSS

***** ---PREPARATIONS-DO K=KTS, LMH-1 DTOZ(K) = DTDIF/(ZHK(K) - ZHK(K+1))CR(K) = -DTOZ(K) * RKH(K)IF(K.LT.LPBL) THEN DO M=MSS,NSPEC RKCT(M,K)=0.ENDDO ELSE RKHZ=RKH(K)*(ZHK(K)-ZHK(K+2))DO M=MSS, NSPEC RKCT(M,K)=RKHZ*CTS(M)*0.5 ENDDO ENDIF ENDDO ---TOP LEVEL RHOK=RHO(KTS) CM(KTS)=DTOZ(KTS)*RKH(KTS)+RHOK DO M=MSS, NSPEC RSS(M, KTS)=-RKCT(M, KTS)*DTOZ(KTS)+SPECIES(M, KTS)*RHOK ENDDO ---INTERMEDIATE LEVELS-DO K=KTS+1, LMH-1 DTOZL=DTOZ(K) CF=-DTOZL*RKH(K-1)/CM(K-1) RHOK=RHO(K) CM(K) = -CR(K-1)*CF+(RKH(K-1)+RKH(K))*DTOZL+RHOKDO M=MSS,NSPEC RSS(M,K) = -RSS(M,K-1)*CF &+(RKCT(M,K-1)-RKCT(M,K))*DTOZL+SPECIES(M,K)*RHOK ENDDO ENDDO -BOTTOM LEVEL DTOZS=DTDIF/(ZHK(LMH)-ZHK(LMH+1)) RKHH=RKH(LMH-1) CF=-DTOZS*RKHH/CM(LMH-1) CMB=CR(LMH-1)*CF RHOK=RHO(LMH) DO M=MSS, NSPEC RKSS=RKHS*CLOW(M) CMSB=-CMB+(RKHH+RKSS)*DTOZS+RHOK RSSB=-RSS(M,LMH-1)*CF+RKCT(M,LMH-1)*DTOZS+SPECIES(M,LMH)*RHOK SPECIES(M,LMH)=(DTOZS*RKSS*SZ0(M)+RSSB)/CMSB ENDDO ---BACKSUBSTITUTION DO K=LMH-1,KTS,-1 RCML=1./CM(K) DO M=MSS, NSPEC SPECIES(M,K)=(-CR(K)*SPECIES(M,K+1)+RSS(M,K))*RCML ENDDO ENDDO END SUBROUTINE VDIFH SUBROUTINE VDIFX(DTDIF,LMH,RKHS,CT

MYJ

&

Summary

- The vortex bogussing technique and other modeling tricks were used to produce landfall simulations of Wilma (2005) with the correct track, intensity, and size.
- Under these conditions, a mesoscale model can produce time series of local wind speeds in good agreement with observations. (Make sure to correct for open exposure when needed.)
- Over water, the boundary layers and surface wind fields produced by the YSU and MYJ schemes are very similar.
- Over land, the MYJ does not appear to recognize further increases in roughness length associated with land surface types. YSU may overestimate the effect.
- The causes for these differences remain to be understood.









